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Quantifying the impact of sulfate geoengineering on mortality from air quality and UV-B exposure

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Highlights:

- Direct, non-climate effects of sulfate injection produce net health risk reduction
- Surface sulfur emission incurs 25 times the exposure from stratospheric injection
- Disbeneficial climate change-driven health effects dominate impacts of injection
- Net impacts of injection harmful despite beneficial photochemical response
- Injection health impacts small relative to risks associated with climate change
Abstract

Sulfate geoengineering is a proposed method to partially counteract the global radiative forcing from accumulated greenhouse gases, potentially mitigating some impacts of climate change. While likely to be effective in slowing increases in average temperatures and extreme precipitation, there are known side-effects and potential unintended consequences which have not been quantified. One such consequence is the direct human health impact. Given the significant uncertainties, we take a sensitivity approach to explore the mechanisms and range of potential impacts. Using a chemistry-transport model, we quantify the steady-state response of three public health risks to 1°C global mean surface cooling. We separate impacts into those which are “radiative forcing-driven”, associated with climate change “reversal” through modification of global radiative forcing, and those “direct impacts” associated uniquely with using sulfate geoengineering to achieve this. We find that the direct (non-radiative forcing driven) impact is a decrease in global mortality of ~13,000 annually. Here the benefits of reduced ozone exposure exceed increases in mortality due to UV and particulate matter, as each unit of injected sulfur incurs 1/25\textsuperscript{th} the particulate matter exposure of a unit of sulfur emitted from surface sources. This reduction is exceeded by radiative forcing-driven health impacts resulting from using sulfate geoengineering to offset 1°C of surface temperature rise. Increased particulate matter formation at these lower temperatures results in ~39,000 mortalities which would have been avoided at higher temperatures. As such we estimate that sulfate geoengineering in 2040 would cause ~26,000 (95% interval: -30,000 to +79,000) early deaths annually relative to the same year without geoengineering, largely due to the loss of health benefits associated with CO\textsubscript{2}-induced warming. These results account only for impacts due to changes in air quality and UV-B flux. They do not account for non-mortality impacts or changes in atmospheric dynamics, and
must be considered in the wider context of other climate change impacts such as heatwave frequency and sea level rise.

Keywords: Geoengineering; air quality; UV exposure; ozone; mortality

1 Introduction

Sulfate geoengineering is one of several possible forms of solar radiation management (SRM), proposed as a method to reduce the net harm resulting from anthropogenic climate change. By promoting the formation of a long-lived stratospheric aerosol layer, a fraction of incoming solar radiation can be scattered back to space before it could be absorbed by the atmosphere, partially offsetting the net anthropogenic radiative forcing. The efficacy of a natural or artificial sulfate layer in reducing global temperature and precipitation has been widely investigated. Early investigations focused on large volcanic eruptions, which are known to produce transient stratospheric aerosol layers (McCormick et al., 1995), while later climate modeling studies explored the possible outcomes of sulfate geoengineering (Rasch et al., 2008).

Although the climate and public health impacts of sulfate geoengineering have been discussed (NAS 1992, Pitari et al., 2014, Effiong et al., 2016), to date there has not been a quantitative evaluation of how global mortality rates might be affected by changes in air quality or UV-B exposure resulting from such a strategy.

Air quality, specifically surface-level concentrations of ozone and fine particulate matter (PM$_{2.5}$), has been linked quantitatively to changes in mortality rates through exposure response functions based on epidemiological studies (Hoek et al., 2014, Jerrett et al., 2010). A similar function has been developed for exposure to UV-B radiation, with the aim of estimating avoided skin cancer incidence due to implementation of the Montreal protocol (Slaper et al., 1996). The
existence of these functions allows the effect of any policy or technology on each of these factors to be calculated and compared in common units. Degraded air quality is estimated to cause ~8% of all global mortality in 2015 (Cohen et al., 2017), and changes to air quality are frequently considered in the context of climate change. A recent study found that mitigation of greenhouse gas emissions from an “unconstrained” scenario down to those in the IPCC RCP4.5 scenario would result in ~1.3 million fewer mortalities per year in 2050 due to both changes in climate and the required changes in emissions (West et al., 2013). This estimation technique has been applied in source-specific impacts evaluations such as for aircraft emissions (e.g. Eastham and Barrett, 2016), but has not yet been applied to an analysis of sulfate geoengineering.

The mechanisms by which an SRM proposal affects these outcomes can be separated into two categories: the “direct” impacts of the method, and the “RF (radiative-forcing)-driven” impacts. Figure 1 gives a conceptual overview of how these categories apply to the impact pathways between stratospheric injection of sulfate aerosol and human health impacts. “Direct impacts”, shown in black, include any effects of the technique which would occur even if there were no effect on the climate. For sulfate geoengineering, an example would be the descent of injected aerosol to the surface. Falling aerosol will add to the existing burden of near-surface fine particulate matter, degrading surface air quality and incurring public health damages in the form of increased respiratory disease mortality rates. This impact would occur regardless of whether the injected aerosol successfully reduced the net radiative imbalance. A second example is the effect of sulfate geoengineering on stratospheric ozone, and the resulting effect on the intensity of surface-level UV-B radiation. Although this has been discussed in the literature in terms of changes in mean intensity (Pitari et al., 2014, Nowack et al., 2016, Xia et al., 2017), the impact on human health has never been quantified.
Figure 1. Influence diagram for impacts of sulfate geoengineering on public health. Only first-order influences are shown here. Minor contributions which may still be significant, such as the direct scattering effect of sulfate aerosols on surface UV-B flux, are not shown for the sake of clarity.

“RF-driven” impacts, shown in Figure 1 in red, include only those which result from the change in radiative forcing achieved by the injected aerosol. Although there are some ways in which the effects of sulfate geoengineering are expected to differ from a simple reversal of climate change (Caldeira et al., 2013), RF-driven impacts are likely to be dominated by the avoided effects of climate change. For example, increasing temperatures associated with climate change are expected to increase ozone concentrations in polluted regions (Fiore et al., 2012). By mitigating future increases in temperature, sulfate geoengineering might reduce total mortality due to ozone exposure relative to the avoided future scenario. Similarly, any potential localized benefits of climate change such as increased crop yields in previously-unproductive regions (Reilly et al., 1994) would also be lost.
The relative contribution of each impact pathway to the total impact of sulfate geoengineering depends on multiple uncertain quantities. Although volcanic events have provided evidence that a stratospheric sulfate layer can provide a negative radiative forcing, the total forcing achieved per unit mass injected varies between studies. For a given target outcome – for example, a 1°C reduction in global average surface temperature – the required rate of sulfate injection will depend on the lifetime and properties of the aerosol layer produced, in addition to the sensitivity of the climate to an increase in stratospheric aerosol optical depth. There are large differences in the estimates of the RF per unit sulfate, differences that depend, in part, on the way sulfates are introduced to the stratosphere (Pierce et al., 2010, Niemeier and Timmreck, 2015). The magnitude of these variables could affect the total impact of sulfate geoengineering, and the contribution of each pathway. Lower RF per unit sulfate means larger direct impacts per unit climate benefit. A world with a low climate sensitivity (the rate of change of temperature with respect to aerosol optical depth, \( \partial T/\partial \tau \)) will require more injected mass to achieve the same temperature reduction target than a world with a high climate sensitivity, but the amount of temperature reduction used in a low-sensitivity world will presumably be correspondingly less. Although temperature-related impacts would be unaffected, impacts directly related to the presence of more stratospheric aerosol, including stratospheric ozone changes, and therefore UV-B exposure, will be greater for the former case than the latter. Although a spot estimate of geoengineering’s impacts on global mortality can be achieved in a single model run, a more nuanced approach is required to understand what the contribution of each pathway is to the total, and how these contributions are affected by uncertainty in input parameters such as climate sensitivity.
We use a global chemistry-transport model (CTM) to compute the response of air quality and population UV-B exposure to sulfate geoengineering at a rate of 1 TgS/yr, isolating the direct and RF-driven impacts using a hybrid modeling approach. Direct impacts of sulfate geoengineering are estimated using offline CTM simulations, in which meteorological fields are specified and no climate response is simulated. RF-driven impacts are estimated by re-running the CTM with perturbed meteorological fields, using a GCM to calculate temperature and precipitation changes resulting from sulfate geoengineering. For each of these simulations, the impact of the relevant pathway is calculated by comparison to a baseline simulation in which no sulfate geoengineering is simulated. Assuming a linear relationship of uncertain slope between stratospheric AOD and temperature change, we apply a Monte-Carlo method to estimate the overall impact of sulfate geoengineering sufficient to achieve a 1°C reduction in global average surface temperature on global mortality due to air quality and UV-B exposure, quantifying the contribution of direct and RF-driven impact pathways to the total.

2 Methods

Air quality and UV-B exposure changes resulting from sulfate geoengineering are calculated using a hybrid modeling approach, combining simulations in a global chemistry-transport model (CTM) with results from a sulfate geoengineering simulation in a global climate model (GCM). For each scenario, impacts are calculated by calculating the difference in results between the output from two CTM simulations.

CTM simulations are performed using prescribed meteorology, so climate feedbacks are decoupled from the atmospheric conditions in the model. RF-driven impacts are simulated by imposing pre-calculated changes in temperature and precipitation directly to the meteorological fields within the CTM. This allows the direct and RF-driven impacts to be isolated, while taking
advantage of the modeling skill of the CTM’s chemical mechanism with respect to simulating changes in air quality and UV-B exposure. The properties of stratospheric aerosol under baseline and geoengineered conditions are also calculated separately, using a dedicated aerosol microphysics model to provide size parameters for each case. The model setup to simulate atmospheric composition in 2040 with and without sulfate geoengineering, and the approach used to disaggregate impact pathways, is described in section 2.1.

These CTM simulations are sufficient to provide a single estimate of the net impact of geoengineering at a rate of 1 TgS/yr on surface air quality and UV-B exposure, in addition to the relative contribution of each direct and RF-driven impact to the total. However, it does not account for uncertainty in the climatological response. By assuming linearity in the relationships between several atmospheric and climatological variables, we convert our estimate of the impact of 1 TgS/yr of aerosol injection into the impact of a specific target climate outcome: offsetting 1°C of global mean surface temperature increase. This method is described in section 2.2.

We also extrapolate the effect of uncertainty in the climate variables to compute the level of uncertainty in the net impact of sulfate geoengineering on air quality and UV-B, holding the target climate outcome constant. The Monte-Carlo method applied to achieve this is described in section 2.3. Finally, we apply epidemiological exposure-response functions to determine the net change in global mortality resulting from achieving this climate outcome, and the relative contribution of direct and RF-driven mechanisms. This is described in section 2.4.
2.1 Atmospheric modeling

Impacts of sulfate geoengineering are computed for a target year of 2040. Atmospheric composition in 2040, with and without sulfate geoengineering and the associated RF-driven impacts, is calculated using the GEOS-Chem atmospheric model.

GEOS-Chem is a global chemistry-transport model (CTM), directly simulating atmospheric chemistry, transport, radiative transfer of UV, emissions, and loss processes. Following the recent implementation of a unified tropospheric-stratospheric chemistry extension, GEOS-Chem uses the same comprehensive chemical mechanism throughout both the troposphere and stratosphere, including an explicit representation of stratospheric aerosols (Eastham et al., 2014). For all CTM simulations we use meteorological fields produced from the NASA GMAO Global Earth Observation System (GEOS-5) for the years 2004-2010. This simulation period is repeated once to yield 14 years of output. The meteorological data is made up of 72 layers from the surface to 0.1 hPa, and is regridded to a horizontal resolution of 4°×5°. Boundary conditions and surface anthropogenic emissions are taken from the RCP 4.5 projection for 2040 (Wise et al., 2009, Clarke et al., 2007, Smith et al., 2006). Initial conditions representative of the future atmosphere are calculated using a prior 14-year spinup simulation, resulting in a total integration time of 28 years. This extended integration time is required to ensure that the model has reached steady state prior to the period of analysis. The effects of geoengineering are calculated by comparing the mean atmospheric state over the final five years between two simulations (e.g. the results of a simulation with 1 TgS/yr injection are compared to a baseline simulation in which no sulfate geoengineering is employed). Surface-level PM$_{2.5}$ and ozone concentrations are retrieved based on the output at the lowest model layer. UV-B exposure is calculated based on the surface-level incident UV radiation fluxes estimated by the Fast-JX
UV radiative transfer and photolysis code embedded in GEOS-Chem, with each wavelength bin weighted according to the SCUP-h action spectrum relevant to UV-induced DNA damage in human skin (de Gruijl and Van der Leun, 1994).

The microphysical properties of the stratospheric aerosol are estimated separately, using the AER 2-D microphysical model (Weisenstein et al., 1997, 2007). Based on the results of these simulations, a log-normal size distribution is estimated and applied to all sulfate-based stratospheric aerosol in the CTM. For baseline conditions, a modal radius of 0.06 µm is used. More details are given in the SI.

We simulate sulfate geoengineering by directly emitting aerosol into the stratosphere. Sulfate is injected at a rate of 1 TgS/yr between 20 and 25 km pressure altitude, from 30°S to 30°N, and over all longitudes. Consistent with the findings of Pierce et al. (2010) and Benduhn et al. (2016), we assume that sulfur is emitted directly as a sulfate aerosol with the target microphysical properties, rather than as SO\(_2\). Based on results from a 1 TgS/yr injection simulation with the microphysical model we impose a log-normal size distribution on the geoengineered aerosol with a modal radius of 0.16 µm, approximately 2.7 times larger (by radius) than under the baseline case. In an initial calibration simulation, we found that this injection rate results in a mean stratospheric aerosol optical depth (AOD) of 0.079, and that the monthly-average stratospheric burden of geoengineering-attributable sulfate varies by less than ±3% over the five years used to calculate the mean atmospheric state. This approach is sufficient to capture direct impacts of sulfate geoengineering in the absence of the climate response. However, capturing RF-driven impacts requires that the climate response to sulfate geoengineering is simulated or imposed within the CTM.
The response of climate variables (e.g. temperature) to sulfate geoengineering is not coupled to atmospheric composition in the CTM. Instead, temperature and precipitation changes are estimated based on GCM results from GeoMIP (Kravitz et al., 2013). In GeoMIP experiment G4, CanESM2 estimated the climate response to a 0.0472 increase in global stratospheric AOD. 

We took the gridded, monthly mean output fields from this simulation and normalized them by the change in AOD to estimate the temperature response per unit change in the stratospheric-average AOD ($\partial T_{3D}/\partial \tau$). The scalar rate of change of global precipitation per unit change in global average surface temperature ($\partial P/\partial T_{sfc}$) was also estimated. These sensitivities are scaled by the mean change in AOD from the calibration simulation to provide an estimate of the change in temperature and precipitation resulting from sulfate geoengineering at a rate of 1 TgS/yr. 

These changes are then applied to the meteorological fields within the CTM to estimate the RF-driven impacts of sulfate geoengineering on air quality and UV-B exposure. Changes in temperature are applied as a 3-D, absolute change in the temperature field, while changes in precipitation are applied as a relative change in the global average precipitation rate. In both cases, seasonal variation is captured by using monthly mean values rather than an annual average. Further information is provided in the SI.
2.2 Calculation of impacts for a fixed injection rate

As described at the start of section 2, we first calculate the total impact of sulfate injection at a rate of 1 TgS/yr, and separate these impacts into direct and RF-driven pathways. To achieve this, we run 6 separate GEOS-Chem simulations, shown in Table 1.

Table 1. Simulation parameters used for each GEOS-Chem model run.

<table>
<thead>
<tr>
<th>Simulation</th>
<th>Sulfate injection</th>
<th>Precipitation adjustment</th>
<th>Temperature adjustment</th>
<th>Chemistry</th>
</tr>
</thead>
<tbody>
<tr>
<td>Baseline (B)</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>Yes</td>
</tr>
<tr>
<td>Calibration (0)</td>
<td>Yes</td>
<td>-</td>
<td>-</td>
<td>Yes</td>
</tr>
<tr>
<td>Central (C)</td>
<td>Yes</td>
<td>Yes</td>
<td>Yes</td>
<td>Yes</td>
</tr>
<tr>
<td>Precipitation sensitivity (S_P)</td>
<td>Yes</td>
<td>-</td>
<td>Yes</td>
<td>Yes</td>
</tr>
<tr>
<td>Temperature sensitivity (S_T)</td>
<td>Yes</td>
<td>Yes</td>
<td>-</td>
<td>Yes</td>
</tr>
<tr>
<td>Inert aerosol (I)</td>
<td>Yes</td>
<td>-</td>
<td>-</td>
<td>-</td>
</tr>
</tbody>
</table>

The net impact of 1 TgS/yr of sulfate geoengineering on air quality and UV-B exposure is estimated by subtracting the results of the baseline simulation (B) from those of the central simulation (C). The exposure resulting from all emissions in a given scenario (e.g. scenario S) can be represented as E(S), such that the change in exposure due to all effects combined is E(C) – E(B). In the central simulation, all impact pathways are simulated together. 1 TgS/yr of aerosol is injected. Air temperatures are decreased relative to the baseline simulation according to the pre-calculated temperature sensitivity field described in section 1.1, scaled by the 0.079 stratospheric AOD estimated from the calibration simulation. Global precipitation is also decreased relative to the baseline. Post-simulation analysis of the simulation C showed a stratospheric AOD increase of 0.075, within 6% of the value used for calibration.
The contribution of direct (non-RF) pathways to the total impact of sulfate geoengineering is estimated using results from the calibration simulation (0). Specifically, the total contribution of direct pathways (all black arrows in Figure 1) to the net impact of sulfate geoengineering on air quality and UV-B exposure is calculated by subtracting the “RF-driven” impact from the total impact, as \([E(C) – E(B)] – [E(C) – E(0)]\). This is equivalent to \(E(0) – E(B)\) and does not account for second-order terms resulting from, for example, the effect of changes in precipitation on the direct pathways. However, these terms are quantified in Appendix B and found to be negligible.

The contribution of each of the RF-driven pathways to the total is isolated by performing two sensitivity simulations. Each is identical to simulation C, but without one of the two climate perturbations. For example, in simulation \(S_T\), sulfate aerosol is injected and global precipitation is reduced, but temperatures are left unperturbed relative to the baseline. The difference in air quality and UV-B impacts between simulation \(S_T\) and simulation C, calculated as (for example) \(E(C) – E(S_T)\), provides an estimate of the contribution of temperature change (an RF-driven impact) to the net impact of sulfate geoengineering. We refer to the contribution of each of the two RF-driven pathways as the “offset warming” and “offset precipitation” impacts, on the basis that these changes are offsetting impacts of climate change. These pathways account for the effect that the geoengineering-attributable change in RF, and therefore the change in climate, has on background air quality and UV-B exposure. Again, cross terms resulting from interaction between temperature- and precipitation-driven impacts are quantified in Appendix B and found to be negligible.

We run one additional simulation to better disaggregate the direct (non-RF-driven) pathways. The contribution of descending injected sulfate aerosol to concentrations of fine
particulate matter at the surface is calculated by performing a separate simulation, without
temperature or precipitation perturbations, in which a chemically unreactive aerosol is injected
(simulation I). This aerosol undergoes the same loss mechanisms as sulfate aerosol. This direct
impact pathway is referred to as the “descending aerosol” pathway. The net impact due to
descending aerosol is simply $E(I)$, as no other aerosol emissions or formation pathways are
included in this simulation.

Any changes in air quality and UV-B exposure observed in simulation 0 which are not
present in this inert simulation are assumed to be the photochemical response of the atmosphere
to the increased stratospheric loading, calculated as $[E(0) – E(B)] – E(I)$. This direct impact
pathway is referred to as the “photochemical” pathway. Specifically, this is the contribution of
photochemical processes to the total impact of sulfate geoengineering after the impacts of RF
changes on background air quality and UV-B exposure have been accounted for.

2.3 Impacts and uncertainty quantification for a fixed target warming offset

The combination of simulations listed in Table 1 provides an estimate of how sulfate
geoengineering at a rate of 1 TgS/yr would impact air quality and UV-B exposure, in addition to
the contribution from each of four direct and RF-driven pathways. Based on the mean climate
sensitivity from CanESM2 and the calculated stratospheric AOD from GEOS-Chem, this is also
the impact of sulfate geoengineering sufficient to offset 1°C of warming. By assuming linearity
in the atmospheric response, these same results can be used to answer a different question: the
contribution of uncertainty in the atmospheric response to both the total impact and the
contributions of each pathway.
We use a Monte-Carlo approach to explore how uncertainty in three climate variables (Table 2) affects the total calculated change in air quality and UV-B exposure, holding constant the target of offsetting 1°C of surface warming. We assume a linear relationship for each of the following pairs of variables, with the slope of each relationship treated as an uncertain variable: between injection rate and stratospheric aerosol burden (the aerosol lifetime); between stratospheric AOD and temperature change (the climate sensitivity); and between temperature change and precipitation change (the hydrological sensitivity). In each Monte-Carlo simulation, an independent draw of these three variables is taken, and the total impact is recalculated by re-weighting the contribution from each of the four pathways.

We assume that a reduction in the hydrological sensitivity will result in a proportional reduction in impacts due to the RF-driven “offset precipitation” pathway. We assume that a reduction in the climate sensitivity will result in a proportional increase in impacts due to the direct impacts, on the assumption that a decreased climate sensitivity implies an increased AOD for the same warming target, and therefore an increased injection rate. Finally, we assume that a decreased aerosol lifetime implies an increase in the direct “descending aerosol” pathway only. This is on the basis that decreased aerosol lifetimes imply an increased injection rate, but the same overall AOD, with no effect on the overall RF achieved.

In each uncertain draw, aerosol lifetime is chosen based on a uniform distribution between 1 and 2.4 years. This range spans most published estimates (Heckendorn et al., 2009, Pierce et al., 2010, Rasch et al., 2008) and includes the lifetime of 2.4 years simulated by GEOS-Chem in the calibration scenario. For the climate and hydrological sensitivity parameters, a value is randomly chosen from a set of four GeoMIP experiment G4 simulations with different
climate models (CanESM2, MIROC-ESM-CHEM, BNU-ESM and GISS-ER-2) (Kravitz et al., 2013). The parameter distributions are shown in detail in Table 2.

Table 2. Uncertain parameters applied in Monte-Carlo simulations when converting simulation output to mortality estimates.

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Distribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Global temperature sensitivity ((-\partial T/\partial \tau)) (K)</td>
<td>Discrete ([-7.2, -7.3, -12, -19])</td>
</tr>
<tr>
<td>Global hydrological sensitivity ((-\partial P/\partial T)) (% K^{-1})</td>
<td>Discrete ([1.7, 2.4, 2.6, 2.9])</td>
</tr>
<tr>
<td>Mean stratospheric aerosol lifetime (years)</td>
<td>Uniform ([1.0 – 2.4])</td>
</tr>
<tr>
<td>Ozone health response ((dM/d\chi)) (% ppbv(^{-1}))</td>
<td>Triangular ([0.100 – 0.104 – 0.107])</td>
</tr>
<tr>
<td>PM(_{2.5}) health response ((dM/d\chi)) (% (µg m(^{-3}))(^{-1}))</td>
<td>Triangular ([0.500 – 1.10 – 1.60])</td>
</tr>
<tr>
<td>UV-B health response dose factor (unitless)</td>
<td>Triangular ([0.2 – 0.6 – 1.0])</td>
</tr>
</tbody>
</table>

This process is described in more detail in the SI, and an assessment of the accuracy of the linearity assumption is performed in Appendices A and B. Changes in second order effects such as climate variability and atmospheric dynamics, which may affect cross-tropopause mass flux and surface-level stagnation, are not modeled but are a clear priority for future work.

One potentially significant feedback which is not considered here is the effect of sulfate geoengineering on cloud formation and properties. The increase in cloud condensation nuclei (CCN) and ice nuclei resulting from the descent of emitted fine aerosol into the upper troposphere could result in increased cirrus cloud formation, an effect which by one estimate
could contribute up to 60% of the net radiative forcing due to sulfate geoengineering (Kuebbeler, Lohmann, and Feichter, 2012). It is also possible that warm cloud formation could be affected by the increase in CCN. Although this is not likely to be significant for this study, in which the maximum injection rate considered is ~5-10% of current anthropogenic sulfur emissions (Smith et al., 2011), scenarios involving higher rates of sulfate geoengineering emissions could result in additional changes to precipitation patterns, intensity, and frequency which could significantly affect surface concentrations of PM$_{2.5}$. Changes in cloud cover would also affect surface UV-B intensity, potentially mitigating the skin cancer damages simulated here.

2.4 Calculation of health impacts

Once the total change in air quality and UV-B exposure for a given uncertain draw has been computed, we convert the simulated changes in population exposure into an estimate of global mortality. The gradients of the exposure response functions (ERFs), which reflect the sensitivity of health outcomes to population exposure, are treated as uncertain variables, with distributions described in Table 2. We use the non-linear Jerrett et al. (2009), Hoek et al. (2013) and Slaper et al. (1996) ERFs for ozone, PM$_{2.5}$ and UV-B exposure respectively. The Hoek ERF was chosen for PM$_{2.5}$ over the more common Krewski et al. (2009) ERF as it is a global meta-analysis of epidemiological studies including Asia, whereas the latter is an in-depth epidemiological study of the USA only. The effect of applying widely-used alternative ERFs for PM$_{2.5}$ such as those of Krewski or Burnett et al. (2014) is quantified, as is the effect of applying concentration thresholds for both PM$_{2.5}$ and ozone.

One thousand draws are performed for all six uncertain variables, using the Sobol pseudo-random sampling sequence to improve convergence. Sensitivity of the results to each input is calculated using the first-order contributions to total variance. This provides an estimate
of the first-order sensitivity indices (Sobol indices), corresponding to the fractional contributions of uncertainty in each input to the total variance in the output (Saltelli et al., 2008).

Two additional scenarios are simulated with alternative assumptions. The first applies region-specific factors to precipitation changes to quantify the relative importance of global and regional precipitation changes in calculating mortality. The second models a hypothetical low-halogen future to account for the relative contributions of anthropogenic halogens in sulfate geoengineering impact calculations (Tilmes et al., 2009, Tilmes et al., 2012, Heckendorn et al., 2009). A full description of the approach used for these simulations is given in the SI.

3 Results

Impacts of implementing sulfate geoengineering sufficient to offset 1°C of surface warming in 2040 are presented below. Direct pathways are discussed first, followed by RF-driven pathways. A summary of the total impacts is provided in the Discussion section. In each case, the calculated change in mortality is the result of the full Monte-Carlo simulation, propagating uncertainty in climate sensitivity, aerosol microphysics, and exposure response. All impacts are calculated for a projected global population in 2040 of 9 billion people (United Nations, 2013)

3.1 Direct impacts

The first of the direct impacts considered is the descent of injected aerosol to the surface, increasing the surface-level concentration of PM$_{2.5}$. We find that this pathway results in an additional 7,400 premature mortalities per year due to degraded air quality (95% interval: 2,300 to 16,000). This implies that injection of aerosol into the stratosphere sufficient to offset 1°C of surface warming would result in a net increase in mortality of the same order of magnitude as
attributable to jet fuel sulfur in 2006 (Barrett et al., 2012), and an order of magnitude lower than the impacts attributable to shipping in 2002 (Corbett et al., 2007). In an additional sensitivity simulation, we simulated continuous emission of an equal mass of aerosol at the surface, distributed according to present-day surface-level sulfur emitters. Per unit mass emitted, we find that surface-level emissions of sulfate result in 25 times greater population exposure to PM$_{2.5}$ than results from emitting the same aerosol into the stratosphere, while achieving a greater radiative forcing offset due to the longer lifetime of stratospheric aerosol.

Direct photochemical changes, excluding the impact of injected aerosol descending to the surface, is net negative, with a mean outcome of -42,000 premature mortalities per year (95% interval: -42,000 to -4,900). This response is dominated by decreased ozone exposure at the surface. Enhanced stratospheric ozone depletion results in reduced ozone mixing ratios in surface-bound stratospheric air masses, while the increased mid-tropospheric flux of UV radiation reduces the photochemical steady-state concentration of ozone throughout the troposphere (Zhang et al., 2014). Changes in the atmospheric dynamics, including the stratosphere-troposphere ozone exchange rate due to dynamical effects of sulfate geoengineering, are not considered but may affect this result (Kirtman et al., 2014). The mean change in global mortality due to reduced ozone exposure in this pathway is -23,000, exceeding the mean increase in skin cancer mortality of 4,100 due to increased UV-B exposure. The reduction in ozone also prompts a small decrease in PM$_{2.5}$, resulting in -1,400 premature mortalities (-2,400 to -520) per year. This suggests that a small depletion in stratospheric ozone may result in a net reduction in global mortality. This is a surprising result, and implies that future increases in stratospheric ozone such as those projected under some climate change scenarios (Li et al., 2009) might be considered as a public health threat. However, this outcome
may be specific to the circumstances of the stratospheric ozone loss, and warrants further research.

Previous studies have shown that the stratospheric ozone loss due to sulfate geoengineering is sensitive to the assumed halogen loading, with one study even finding a reversal of sign (Tilmes et al., 2009, Tilmes et al., 2012, Heckendorn et al., 2009). We simulate an alternative scenario which corresponds to the theoretical minimum atmospheric halogen loading. In this scenario all anthropogenic halogen emissions are set to zero, as are the initial concentrations for all long-lived anthropogenic halogen gases (see SI for details). We find that total ozone column depletion is reduced by 31% relative to the scenario with RCP 4.5 halogen emissions, resulting in 2,500 fewer premature mortalities due to skin cancer, and 4,800 fewer due to PM$_{2.5}$ exposure. These benefits are exceeded by the increased ozone exposure in this scenario, resulting in 7,600 additional mortalities. The net result is that the reduction in global mortality due to direct photochemical impacts alone is smaller in magnitude by 3.6% under a low-halogen scenario, relative to the baseline scenario. This again suggests that a relative increase in ozone concentrations may have a net public health disbenefit, considering only air quality and UV-B exposure.

Considering only direct impact pathways, sulfate geoengineering sufficient to offset 1°C of surface warming results in a net benefit, with a global change of -13,000 premature mortalities per year (sum of central estimates). Although we find 7,400 (2,300 to 16,000) additional premature mortalities due to direct population exposure to injected aerosol, this is counteracted by -20,000 (-42,000 to -4,900) premature mortalities due to photochemical impacts resulting from the increased sulfur loading of the stratospheric aerosol layer.
3.2 RF-driven impacts

The calculated RF-driven impacts of sulfate geoengineering on air quality are consistent with prior literature examining the related problem of the response of air quality to CO$_2$-driven warming. Reduced temperatures relative to the projected future scenario result in enhanced partitioning of HNO$_3$ from background emissions into nitrate aerosol, and therefore an increase in surface PM$_{2.5}$. We find that this dominates other PM$_{2.5}$ formation mechanisms which reduce in response to cooler surface temperatures, such as production of biogenic aerosols. The result is that, by offsetting 1°C of surface warming from climate change, sulfate geoengineering results in an additional 69,000 premature mortalities annually (41,000 to 95,000). This increase is accompanied by a significant decrease in premature mortality due to the avoided effect of global warming on ozone. Ozone concentrations in polluted regions decrease with temperature as photochemical production is slowed, such that sulfate geoengineering results in -43,000 (-67,000 to -19,000) premature mortalities per year due to ozone exposure relative to the avoided future.

The effect of temperature change on UV-B exposure is negligible.

The other RF-driven impact of sulfate geoengineering is lower overall precipitation rates, offsetting some of the increased precipitation projected to result from climate change. Decreased precipitation results in longer lifetimes for PM$_{2.5}$ and therefore in increased PM$_{2.5}$ exposure globally. The total RF-driven impact of changes in precipitation from sulfate geoengineering is an additional 14,000 (7,100 to 21,000) premature mortalities per year, with negligible effects on ozone and UV-B exposure.

This approach assumes that precipitation will be uniformly affected across all locations, and all aggregated impacts in the following sections are calculated on this assumption. However, sulfate geoengineering is likely to reduce precipitation by a greater proportion in some regions...
than in others (Kravitz et al., 2014). We run an additional, sensitivity simulation in which precipitation rate modifications are derived and applied on a local basis rather than a global basis, using a separate factor derived from the GeoMIP G4 CanESM2 simulation for each of 21 climatologically-distinct regions (see SI for details). In this scenario, the impacts of precipitation are increased by 15%. The increase occurs almost exclusively in Asia and Eastern Europe. Here the relative reduction in precipitation is 1.5 and 4.5 times the global average, respectively, resulting in longer lifetimes for PM$_{2.5}$ as washout is decreased. However, whether using global or regional precipitation adjustments, impacts due to temperature change remain dominant factor in RF-driven mortality impact pathways of sulfate geoengineering.

The net effect of RF-driven impact pathways on global air-quality and UV-B exposure is a net increase in mortality, reflecting the loss of climate change-driven air quality benefits associated with increasing temperature and precipitation. We find a combined central estimate of 39,000 additional mortalities per year due to this offsetting. This total is made up of +26,000 (-12,000 to +63,000) premature mortalities due to avoided temperature change, and +13,000 (+6,600 to +20,000) due to precipitation reduction.

These results are sensitive to the modeled impact of climate change on surface air quality. We find that, for both RF-driven pathways, sulfate geoengineering offsets climate change-related increases in ozone and decreases in PM$_{2.5}$, with the magnitude of mortality impacts from the latter change exceeding those from the former. While increases of ozone under climate change are widely reported in the literature, the sign of the impact of climate change on surface PM$_{2.5}$ concentrations is uncertain (Fiore et al., 2012). The net outcome of RF-driven impacts is likely to change as our understanding of the impacts of climate change is further refined.
4 Discussion

In total, and considering only the effects on air quality and UV-B exposure, sulfate geoengineering sufficient to offset 1°C of warming results in +26,000 premature mortalities annually (95% confidence interval of -30,000 to +79,000). Figure 2 shows a graphical breakdown of incurred mortalities by pathway and by exposure type, with numerical values shown in Table 3. This total is made up of 39,000 additional mortalities due to RF-driven pathways, partially offset by 26,000 prevented mortalities due to direct pathways, as outlined in Figure 1. In 17% of cases, mortality reductions due to decreased ozone exposure exceed the combined global mortality impacts of increased PM$_{2.5}$ and UV-B exposure, resulting in a net decrease in global mortality due to sulfate geoengineering. Overall, surface air quality and skin cancer impacts are dominated by increases in mortality due to RF-driven pathways, whereas the direct impact pathways of sulfate geoengineering are net beneficial by these metrics.
Figure 2. Annual premature mortality impacts resulting from sulfate geoengineering sufficient to offset 1°C of surface warming. Impacts are separated by pathway, based on Figure 1, and by exposure type. The left panel shows the contributions to each pathway’s total impact, separated by exposure type. The right hand sub-plot shows how each impact pathway contributes to the total. “Descending inj. mass” corresponds to direct exposure of the population to injected aerosol mass as it descends to the surface. “Photochem. effects” corresponds to photochemical changes resulting from the increased aerosol optical depth and surface area, including induced changes in stratospheric ozone columns. “Offset warming” corresponds to temperature change, “Offset precip.” to reductions in precipitation. Solid bars show the mean value of Monte Carlo simulation outcomes (n = 1,000). Error bars show the 2.5 and 97.5th percentile values.

For all four atmospheric mechanisms, mortality due to surface ozone exposure is consistently decreased by sulfate geoengineering, whereas mortality due to PM$_{2.5}$ exposure varies in sign. Increases in nitrate aerosol due to reduced surface warming result, on average, in greater
health impacts than the benefits associated with reduced ozone. UV-B exposure is only significantly affected by direct photochemical effects of sulfate geoengineering, but the contribution of changes in UV-B exposure to the overall impact of sulfate geoengineering is an order of magnitude smaller than the contributions of changes in ozone or PM$_{2.5}$ exposure.

When considering only the direct photochemical consequences of sulfate geoengineering, the total skin cancer mortality increase is exceeded by the ozone mortality decrease in all uncertain variable draws. This counterintuitively implies that limited stratospheric ozone destruction may be of net benefit in terms of premature mortality and human lifespan, and that reduction of anthropogenic halogen emissions may increase rather than reduce health impacts due to sulfate geoengineering. However, this does not take into account non-mortality outcomes of exposure to UV-B such as cataract formation and non-melanoma skin cancer, which is less fatal but several orders of magnitude more common than melanoma skin cancer (Guy et al., 2015, Slaper et al., 1996).

Uncertainty in the ERFs for PM$_{2.5}$ and ozone have the greatest first-order effect on overall variance in the global mortality impact of sulfate geoengineering, contributing 44% and 50% of the total variance in the result based on the calculated sensitivity indices. The first-order effects of uncertainty in climate response are an order of magnitude smaller, with the greatest contribution being 2.5% for temperature sensitivity with respect to optical depth. Uncertainty in the UV-B exposure response function, sensitivity of precipitation to temperature and uncertainty in aerosol lifetime each contribute 1% or less to overall uncertainty in the result. When calculating mortality due to PM$_{2.5}$ and ozone individually, ERF uncertainty remains the greatest contributor to overall variance, followed by uncertainty in the temperature sensitivity to optical depth. However, this ordering is reversed for mortalities due to UV-B exposure. Furthermore,
application of an alternative ERF developed for global studies by Burnett et al (2014) results in mortality due to geoengineering-attributable PM$_{2.5}$ falling by 22%. Mortalities calculated using several other ERFs are shown in the SI.

Table 3. Annual premature mortality impacts resulting from sulfate geoengineering sufficient to offset 1°C of surface warming. Mean outcomes are in bold, 95% intervals are shown in square brackets (N = 1,000). The 95% interval is calculated as the 2.5 and 97.5th percentile values of the Monte Carlo simulation outcomes.

<table>
<thead>
<tr>
<th></th>
<th>Direct impacts</th>
<th>RF-driven impacts</th>
<th>All mechanisms</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Descending injection mass</td>
<td>Photochemical effects</td>
<td>Offset warming</td>
</tr>
<tr>
<td>Surface ozone</td>
<td></td>
<td>-23,000</td>
<td>-43,000</td>
</tr>
<tr>
<td>(23,000 : -66,000)</td>
<td></td>
<td>(-67,000 : -19,000)</td>
<td>(-1,100 : -260)</td>
</tr>
<tr>
<td>PM$_{2.5}$</td>
<td></td>
<td>-1,400</td>
<td>69,000</td>
</tr>
<tr>
<td>(2,300 : 16,000)</td>
<td></td>
<td>(-2,400 : -520)</td>
<td>(41,000 : 95,000)</td>
</tr>
<tr>
<td>UV-B</td>
<td></td>
<td>4,100</td>
<td>400</td>
</tr>
<tr>
<td>(1,300 : 8,200)</td>
<td></td>
<td>(200 : 610)</td>
<td>(-40 : -10)</td>
</tr>
<tr>
<td>All causes</td>
<td>7,400</td>
<td>-20,000</td>
<td>26,000</td>
</tr>
<tr>
<td>(2,300 : 16,000)</td>
<td>(-42,000 : -4,900)</td>
<td>(-12,000 : 63,000)</td>
<td>(6,600 : 20,000)</td>
</tr>
</tbody>
</table>

All simulations were performed at a relatively coarse horizontal resolution (4°×5°). A 2013 study indicated that while surface ozone exposure is insensitive to grid resolution, use of coarse horizontal resolution when calculating outcomes could result in mortality due to PM$_{2.5}$ exposure being biased low by 30-40%. This is due to the covariance of peaks in PM$_{2.5}$ concentration and population centers, which is not reflected at coarse resolution (Punger et al., 2013). However, changes in PM$_{2.5}$ due to sulfate geoengineering are diffuse compared to modern anthropogenic PM$_{2.5}$, and this covariance is therefore likely to be reduced. These simulations also do not take into account the possible response of cloudiness to the increase in cloud condensation nuclei which could result from sulfate geoengineering, due to the descent of
emitted fine aerosol into the upper troposphere. In addition to potentially affecting the total UV-B reaching the surface and the net RF associated with geoengineering (Kuebbler et al., 2012), changes in cloudiness through this mechanism could affect surface precipitation and thereby PM$_{2.5}$ concentrations. Although outside the scope of this work, we consider assessment of the response of cloudiness to be a priority for future research on surface-level impacts of geoengineering.

These results must be weighed against the risks of climate change which sulfate geoengineering seeks to mitigate, and the magnitude of current and future health impacts due to degraded air quality. A study of the 2015 global burden of disease found that 4 million deaths annually are attributable to degraded air quality, while air quality co-benefits of greenhouse gas mitigation (including changes in precursor emissions) have been estimated at ~1.3 million fewer mortalities per year in 2050 (Cohen et al., 2017; West et al., 2013). We find that the total air quality and skin cancer related impacts of sulfate geoengineering sufficient to induce a 1°C decrease in surface temperature are +26,000 (95% CI: -30,000 to +79,000) premature mortalities per year. Normalizing by total population in 2040, this is equivalent to a change of +0.3 early deaths per 100,000 population. For context, this can be compared to projected direct health impacts of rising surface temperatures. A study of temperature-related mortality under a “business as usual” (BAU) climate change scenario projected that a 3°C increase in average surface temperature would result in an additional 63,000 mortalities per year in the US alone, corresponding to +20 deaths per 100,000 population (Deschênes et al., 2011). These changes are dominated by increased vulnerability during extreme cold and extreme heat events, resulting in greater changes at higher baseline temperatures. Another study found that aggregate economic impacts of temperature increases are approximately linear in temperature, and that BAU climate
change is estimated to reduce global average incomes by 23% within the next 80 years (Burke et al., 2015). These consequences of climate change must be weighed against the risks and benefits of sulfate geoengineering, including (but not limited to) the impacts on air quality and UV exposure explored in this study, which are relatively small and of uncertain sign.

5 Conclusions

We identify several mechanisms by which sulfate geoengineering may cause changes in air quality and UV-B exposure, and we provide the first quantitative estimates of the impact of sulfate geoengineering on global mortality rates from these causes. When sulfate geoengineering is used to offset 1°C of temperature rise (or create 1°C cooling) we find that RF-driven impacts, associated with offsetting the effects of climate change, result in a net increase in mortality, while other (“direct”) impacts result in a net decrease. The net effect is an increase of 26,000 additional premature mortalities per year (95% interval: -30,000 to +79,000), although the overall sign of the impact is uncertain. We find an 83% chance of a net increase in global mortality due to air quality and UV-B exposure, with uncertainty in the exposure response functions providing the greatest contribution to total uncertainty in the result.

Of the direct impact pathways considered, descent of injected sulfate aerosol from the stratosphere is found to be a minor contributor to the overall impact of sulfate geoengineering. The contribution of descending, injected aerosol to surface PM$_{2.5}$ causes 7,400 additional premature mortalities per year, compared to a decrease of 20,000 premature mortalities per year resulting from the direct photochemical effects of sulfate geoengineering. This is made up of 4,100 additional skin cancer mortalities offset by 23,000 averted premature mortalities due to decreased ozone exposure. By contrast, RF-driven impacts of sulfate geoengineering are found to result in a net increase in mortality relative to the avoided future scenario. By offsetting 1°C of
atmospheric warming, greater concentrations of PM$_{2.5}$ are formed from existing emissions, resulting in an additional 69,000 premature mortalities per year. The reduction in radiative forcing also offsets some of the anticipated increase in precipitation associated with climate change, with longer aerosol lifetimes incurring an additional 14,000 premature mortalities per year. These effects are partially offset by 44,000 avoided mortalities per year from RF-driven changes in ozone exposure. The specific magnitudes depend on the amount of warming which is being offset. The impacts of larger or smaller amounts of can be approximated by scaling the warming to our 1°C value.

This analysis does not account for ecological and climate feedback effects related to increased CO$_2$, possible induced or suppressed cloudiness, or public health impacts beyond changes in mortality due to air quality and UV-B exposure. Deschênes et al. (2011) found that, under a business-as-usual scenario with 3°C of warming in 2070-2099, the direct impact of increased temperatures due to climate change would be 63,000 premature mortalities per year from extreme temperatures in the United States alone. Burke et al. (2015) estimated that aggregate economic impacts of climate change will reduce global average incomes by 23% in the same period. Although beyond the scope of this paper, weighing the broader effects of mitigating climate change against the air quality and UV-B impacts computed here would provide a more complete understanding of the net benefits and risks of sulfate geoengineering.

**Appendix A: Assessment of response linearity**

Four additional simulations are conducted to test the validity of the assumption that impacts will scale linearly with input. For impacts due to changes in temperature and precipitation, we simulate perturbations which are 5 times smaller than the CanESM2 output and
8.6 times smaller than the “full” perturbations corresponding to a 0.98 K cooling. A 0.5 TgS/yr injection rate, resulting in a 0.040 increase in stratospheric optical depth, was simulated to determine impact linearity with respect to these quantities in isolation from meteorological feedbacks. A full list is given in Table 4. Second order effects due to effect interaction (e.g. between precipitation impacts and injected sulfates) are addressed in Appendix B. The output metric shown is the total mortalities as calculated without accounting for uncertainty in climate or exposure response variables.
Table 4. Perturbation parameters used in simulations to establish response linearity.

<table>
<thead>
<tr>
<th>Input</th>
<th>Small perturbation</th>
<th>Large perturbation</th>
</tr>
</thead>
<tbody>
<tr>
<td>Injection rate (TgS/yr)</td>
<td>0.5</td>
<td>1.0</td>
</tr>
<tr>
<td>Optical depth (-)</td>
<td>0.040</td>
<td>0.079</td>
</tr>
<tr>
<td>Offset warming (K)</td>
<td>0.12</td>
<td>0.98</td>
</tr>
<tr>
<td>Offset precipitation (%)</td>
<td>0.28%</td>
<td>2.4%</td>
</tr>
</tbody>
</table>

The total mortalities calculated for each perturbation, broken down into those resulting from exposure to PM$_{2.5}$, ozone and UV-B, are shown in Figure 3. Interpolation between zero and the ‘full-scale’ perturbation shows a good agreement with the results of the smaller test perturbation simulations. The exception to this is in the case of the response to an increase in stratospheric aerosol optical depth $\Delta \tau$. For a $\Delta \tau$ of 0.040, the change in mortality due to skin cancer is 21% greater than would be calculated by interpolation from the impact of a $\Delta \tau$ of 0.079, and the ozone reduction is 13% greater. The effect on PM$_{2.5}$ exposure is negligible. This is likely to be due to saturation, as reaction rates become limited by factors other than surface area density of aerosol.
Figure 3. Response linearity with respect to each of the assessed mortality mechanisms. The dashed line represents the linear sensitivity used in each case when scaling calculated exposures for the purposes of uncertainty quantification.

Appendix B: Second-order sensitivities

Four additional simulations are conducted in which the inputs are combined to determine the effect of second order terms on the response. In the first three simulations, combinations of two parameters (temperature change, precipitation change, and injection rate) are changed simultaneously. In the final simulation, all three are modified together. For these simulations, the effect of descending aerosol and the photochemical effect of an increase in stratospheric optical depth are not separated.
The results of these simulations are shown in Figure 4. In each panel, the left-hand bar shows the total mortalities as calculated by linearly adding the exposure calculated by individual simulations, whereas the right-hand bar shows total mortalities as calculated using a single simulation in which the perturbations are simulated together. These estimates do not include uncertainty in climate variables. Inclusion of second-order effects changes the total calculated number of mortalities by less than ±1%, suggesting that interaction between the three factors is not significant. However, this does not address possible meteorological feedbacks such as changes in cloud cover or ventilation.

Figure 4. Comparison of mortalities estimated by linear combination of calculated exposures from several perturbation simulations (left) and by direct simulation of multiple perturbations together (right). All totals agree to within ±1%. Uncertainty in climate variables is not included in these estimates.
Competing Interests

The authors have no competing interests to declare.

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Acknowledgments, Samples, and Data

All authors designed the research and wrote the paper. DKW undertook the microphysical modeling. SDE undertook the GEOS-Chem modeling. The authors have no financial conflicts of interest to declare.

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References


Intergovernmental Panel on Climate Change, edited by T. F. Stocker, D. Qin, G.-K.

Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, and P. M.


Lim, S.S., Vos, T., Flaxman, A.D., Danaei, G., Shibuya, K., Adair-Rohani, H., Amann, M.,
Balakrishnan, K., Balmes, J., Barker-Collo, S., Baxter, A., Bell, M.L., Blore, J.D., Blyth,
F., Bonner, C., Borges, G., Bourne, R., Boussinesq, M., Brauer, M., Brooks, P., Bruce,
N.G., Brunekreef, B., Bryan-Hancock, C., Bucello, C., Buchbinder, R., Bull, F., Burnett,
R.T., Byers, T.E., Calabria, B., Carapetis, J., Carnahan, E., Chafe, Z., Charlson, F., Chen,
S., Darling, S., Davis, A., Degenhardt, L., Dentener, F., Des Jarlais, D.C., Devries, K.,
Erwin, P.J., Fahimi, S., Falder, G., Farzadfar, F., Ferrari, A., Finucane, M.M., Flaxman,
S., Fowkes, F.G.R., Freedman, G., Freeman, M.K., Gakidou, E., Ghosh, S., Giovannucci,
E., Gmel, G., Graham, K., Grainger, R., Grant, B., Gunnell, D., Gutierrez, H.R., Hall, W.,
Ibeanusi, S.E., Jacklyn, G.L., Jasrasaria, R., Jonas, J.B., Kan, H., Kanis, J. A., Kassebaum,
N., Kawakami, N., Khang, Y.-H., Khatibzadeh, S., Khoo, J.-P., Kok, C., Laden, F.,
Lalloo, R., Lan, Q., Lathlean, T., Leasher, J.L., Leigh, J., Li, Y., Lin, J.K., Lipshultz,
S.E., London, S., Lozano, R., Lu, Y., Mak, J., Malekzadeh, R., Mallinger, L., Marcenes,
W., March, L., Marks, R., Martin, R., McGale, P., McGrath, J., Mehta, S., Mensah, G. A.,
Merriman, T.R., Micha, R., Michaud, C., Mishra, V., Mohd Hanafiah, K., Mokdad, A. A.,
Morawska, L., Mozaffarian, D., Murphy, T., Naghavi, M., Neal, B., Nelson, P.K., Nolla,
J.M., Norman, R., Olives, C., Omer, S.B., Orchard, J., Osborne, R., Ostro, B., Page, A.,
Pandey, K.D., Parry, C.D.H., Passmore, E., Patra, J., Pearce, N., Pelizzari, P.M., Petzold,
Rehm, J.T., Ritz, B., Rivara, F.P., Roberts, T., Robinson, C., Rodriguez-Portales, J. a.,
Romieu, I., Room, R., Rosenfeld, L.C., Roy, A., Rushton, L., Salomon, J. a., Sampson,
U., Sanchez-Riera, L., Sanman, E., Sapkota, A., Seedat, S., Shi, P., Shield, K., Shivakoti,
R., Singh, G.M., Sleet, D. a., Smith, E., Smith, K.R., Stapelberg, N.J.C., Steenland, K.,
Stöckl, H., Stovner, L.J., Straif, K., Straney, L., Thurston, G.D., Tran, J.H., Van
Dingenen, R., van Donkelaar, A., Veerman, J.L., Vijayakumar, L., Weintraub, R.,
Weissman, M.M., White, R. a., Whiteford, H., Wiersma, S.T., Wilkinson, J.D., Williams,
H.C., Williams, W., Wilson, N., Woolf, A.D., Yip, P., Zielinski, J.M., Lopez, A.D.,
risk assessment of burden of disease and injury attributable to 67 risk factors and risk
factor clusters in 21 regions, 1990-2010: a systematic analysis for the Global Burden of


Niemeier, U., and Timmreck, C. (2015), What is the limit of climate engineering by stratospheric

Nowack, P. J., N. L. Abraham, P. Braesicke, and J. A. Pyle (2016), Stratospheric ozone changes
under solar geoengineering: implications for UV exposure and air quality, Atmos. Chem.

Outdoor air pollution: Assessing the environmental burden of disease at national and


Dalsoren, S., Faluvegi, G., Folberth, G., Horowitz, L.W., Nagashima, T., Naik, V.,
Rumbold, S., Skeie, R., Sudo, K., Takemura, T., Bergmann, D., Cameron-Smith, P.,
Cionni, I., Doherty, R.M., Eyring, V., Josse, B., MacKenzie, I. a., Plummer, D., Righi,
mortality due to anthropogenic outdoor air pollution and the contribution of past climate
of ozone depletion and skin cancer incidence to examine the Vienna Convention
Smith, S.J., Aardenne, J. van, Klimont, Z., Andres, R.J., Volke, A., and Delgado Arias, S.,
1101–1116.
Smith, S. J., and T. M. L. Wigley (2006), Multi-Gas Forcing Stabilization with MiniCAM,
Energy J., (Special Issue #3), 373–391.
Tilmes, S., R. R. Garcia, D. E. Kinnison, A. Gettelman, and P. J. Rasch (2009), Impact of
geoengineered aerosols on the troposphere and stratosphere, J. Geophys. Res., 114(D12),
D12305.
Tilmes, S., D. E. Kinnison, R. R. Garcia, R. Salawitch, T. Canty, J. Lee-Taylor, S. Madronich,
and K. Chance (2012), Impact of very short-lived halogens on stratospheric ozone
abundance and UV radiation in a geo-engineered atmosphere, Atmos. Chem. Phys., 12,
10945–10955.


Wise, M., K. Calvin, A. Thomson, L. Clarke, B. Bond-Lamberty, R. Sands, S. J. Smith, A.
Janetos, and J. Edmonds (2009), Implications of limiting CO2 concentrations for land use
and energy, Science, 324(5931), 1183–1186.

goengineering on tropospheric ozone. Atmospheric Chemistry and Physics, 17(19),
11913–11928.

Zhang, H., S. Wu, Y. Huang, and Y. Wang (2014), Effects of stratospheric ozone recovery on
photochemistry and ozone air quality in the troposphere, Atmos. Chem. Phys., 14, 4079–
4086.